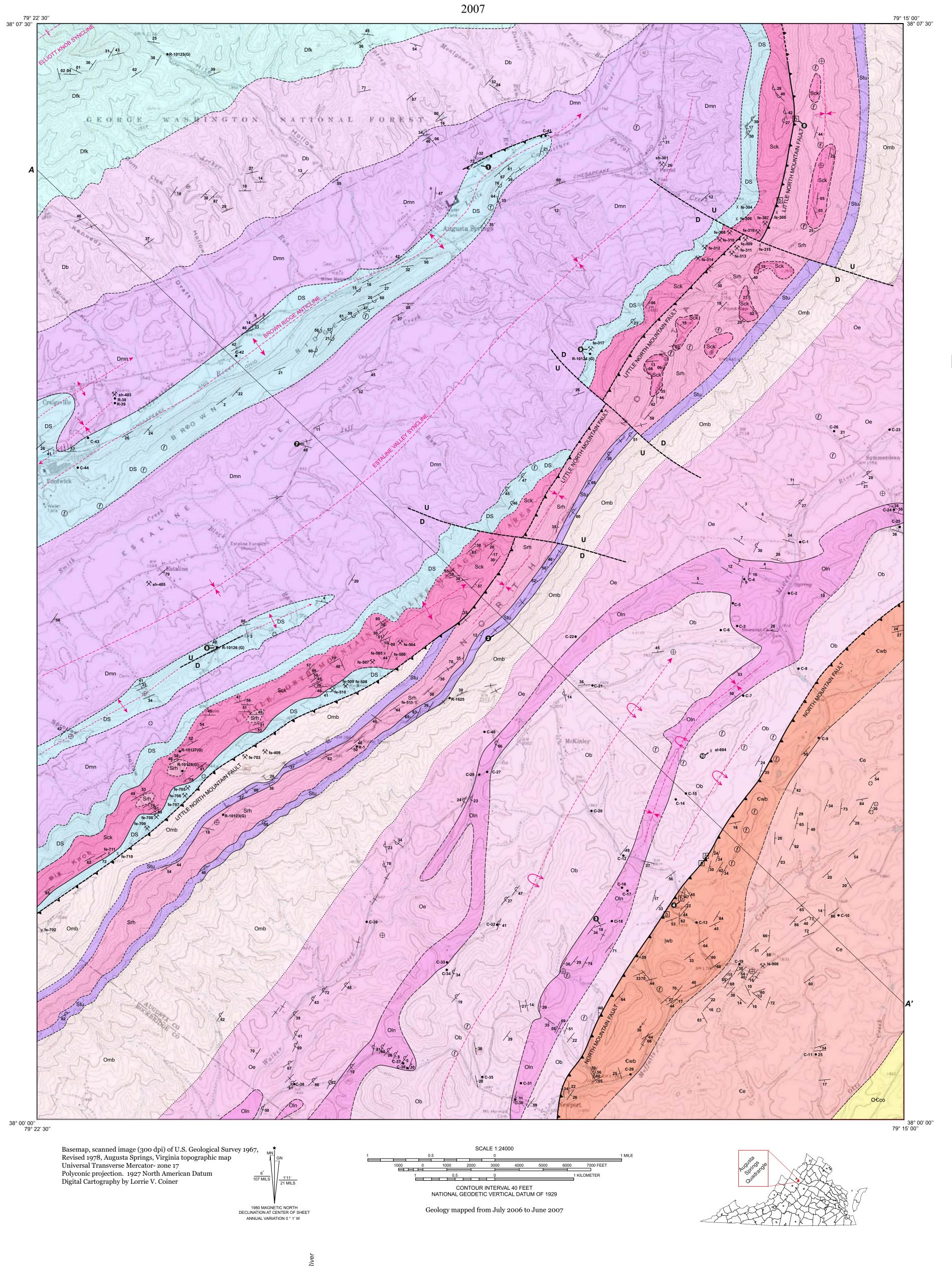
ELEVATION (ft)

BEDROCK GEOLOGIC MAP OF THE AUGUSTA SPRINGS QUADRANGLE, VIRGINIA

LORRIE V. COINER AND GERALD P. WILKES



LITTLE NORTH

ESTALINE VALLEY

SYNCLINE

BROWN RIDGE

ANTICLINE

MOUNTAIN FAULT

DESCRIPTION OF MAP UNITS

Foreknobs Formation - Green and greenish-gray, thin- to thick-bedded, fossiliferous (most notably large crinoid stems) quartz sandstone and shale and minor quartz-pebble conglomerate. Only the lowest part of the Foreknobs is exposed in the quadrangle.

Brallier Formation - Olive-gray, thin-bedded, micaceous, sparsely fossiliferous siltstone, shale, and thin lithic sandstone. Thickness: 1500 to 2200 feet. Conformable with the overlying Foreknobs Formation.

Millboro Shale and Needmore Formation, undivided - Millboro Shale: black, fissile shale, with thin bentonite beds. Near the base is an interval of dark-gray, aphanic, thin-bedded limestone (Purcell Member?). Needmore Formation: olive-gray, weathered, fossiliferous shale, with thin bentonite beds. Composite thickness 800 to 1200 feet. Conformable with the overlying Brallier Formation.

Devonian and Silurian rocks, undivided - Ridgeley Sandstone: medium-gray, fine- to coarse-grained, medium- to thick-bedded, calcareously cemented, fossiliferous (most notably "Spirifer" brachiopods) quartz sandstone. Helderburg Group, including Licking Creek Limestone: medium-gray, sandy, thick-bedded limestone with abundant white chert (Figure 1). New Creek Limestone: gray to pink, coarse-grained, medium-to thick-bedded, crinoidal limestone. Keyser Formation: medium-gray, coarse-grained, thick-bedded, nodular- weathering limestone. Tonoloway Limestone: dark-gray, fine-grained, thin-bedded, argillaceous limestone. Thickness of Devonian-Silurian unit is from 255 to 820 feet. Unconformable with the overlying Needmore Formation.

Cayugan Group and Keefer Sandstone - Cayugan Group includes rock units that have interfingering facies relationships. Wills Creek Formation: yellow-weathering, silty shale. Williamsport Sandstone: greenish-gray, medium- to coarse-grained, medium- to thick-bedded, indurated quartz sandstone, with minor siltstone beds. McKenzie Formation: calcareous, crossbedded sandstone, interbedded with medium-to dark-gray, medium- to thick-bedded, knobby-weathering limestone. Map unit also includes the Bloomsburg Formation: maroon and light-gray, fine-grained, medium-bedded quartz sandstone (locally referred to as "Leopard Rock" due to mottling on bedding surfaces), red and tan mudstone. Keefer Sandstone: white to light-gray, medium- to coarse-grained, thick-bedded, ledge-forming quartz sandstone, quartzitic in places. Thickness of the Cayugan Group and Keefer Sandstone is 150 to 800 feet. Conformable with the overlying Tonoloway Limestone.

Rose Hill Formation - Maroon, medium-grained, thin- to thick-bedded, hematitic, indurated quartz sandstone, and red, green, and gray shale and siltstone. Thickness: 200 to 800 feet. Conformable with the overlying Keefer Sandstone.

Tuscarora Formation - White to light-gray, medium- to coarse-grained, thick-bedded, ledge-forming quartz sandstone and quartzite (Figure 2), with quartz-pebble conglomerate ("jelly beans") in lower half of unit. Thickness: 50 to 260 feet. Conformable with the overlying Rose Hill Formation.

Martinsburg Formation - Upper unit: light- to medium-gray, fine-grained, thin-

Lincolnshire Limestone and New Market Limestone - Lincolnshire Limestone:

black chert, beds of white massive chert near top of unit. Some interbeds of light- to

quartz sandstone. Silty laminae in the Conococheague stand out in relief on weathered

surfaces. Thickness: 2000 to 2500 feet. Conformable with the overlying Beekmantown

MAP SYMBOLS

For all contact, fault and fold symbols: lines are solid where the location is exact, long-dashed where the location is approximate, short-dashed where the location is

inferred; dotted where the location is concealed. Queries added where identity or existence may be questionable. For geologic observation symbols, observation sites

bedded sandstone, siltstone, and shale, distinctive Orthorhynchula zone near top of unit. Lower unit: light-gray, argillaceous, massive- to thin-bedded limestone. Up to 1600 feet thick. Unconformable with the overlying Tuscarora Sandstone.

De Edinburg Formation - Liberty Hall facies: black to dark-gray, aphanic, medium- to thick-bedded limestone, with thin black shale interbeds. Lantz Mill facies: black, medium- to coarse-grained, thick-bedded, nodular weathering limestone. Thickness: 800 to 1250 feet. Graded contact with the overlying Martinsburg Formation.

dark-gray, fine- to coarse-grained, thick-bedded limestone, with black chert (Figure 3).

New Market Limestone: light-gray, aphanic, thick-bedded limestone. Composite thickness 100 to 435 feet. Conformable with the overlying Edinburg Formation.

Beekmantown Formation - Light- to medium-gray, fine-grained, thick-bedded dolostone. Distinctive "butcher-block" jointing on weathered bedding surfaces, bedded

medium-gray, medium-bedded limestone. Thickness: 1500 to 2000 feet. Unconformable with the overlying New Market Limestone.

Conococheague Formation - Medium-gray, fine-grained, thin- to thick-bedded limestone. Some interbeds of light-gray, fine- to medium-grained, medium-bedded dolostone, and calcareously cemented, fine- to medium-grained, medium-bedded

Elbrook Formation - Light- to dark-gray, fine- to medium-grained, thin-bedded dolostone and limestone, and pink and green shale. Thickness: 1325 to 2300 feet. Conformable with the overlying Conococheague Formation.

Waynesboro Formation - Maroon, green, and gray shale, interbedded with medium-gray, fine-grained limestone with minor black chert nodules, and light- to dark-gray, fine- to coarse-grained, calcareous dolostone. Only the upper part of the Waynesboro is exposed in the quadrangle. Conformable with the overlying Elbrook Formation.

are centered on the strike bar or at the intersection point of multiple symbols.

Thrust - sawteeth on upthrown block

Normal - U - upthrown block, D - downthrown block

Folds - showing direction of plunge where appropriate

----- Stratigraphic Contacts

Fault Contacts

Overturned syncline

Geologic Observations

Strike of vertical beds

Horizontal bedding

Strike and dip of inclined beds

Strike and dip of overturned beds

Strike and dip of inclined foliation

Strike and dip of deformed bedding

lineation (pencil cleavage)

Strike of vertical foliation

Interpretive cross-section:

1. No vertical exaggeration

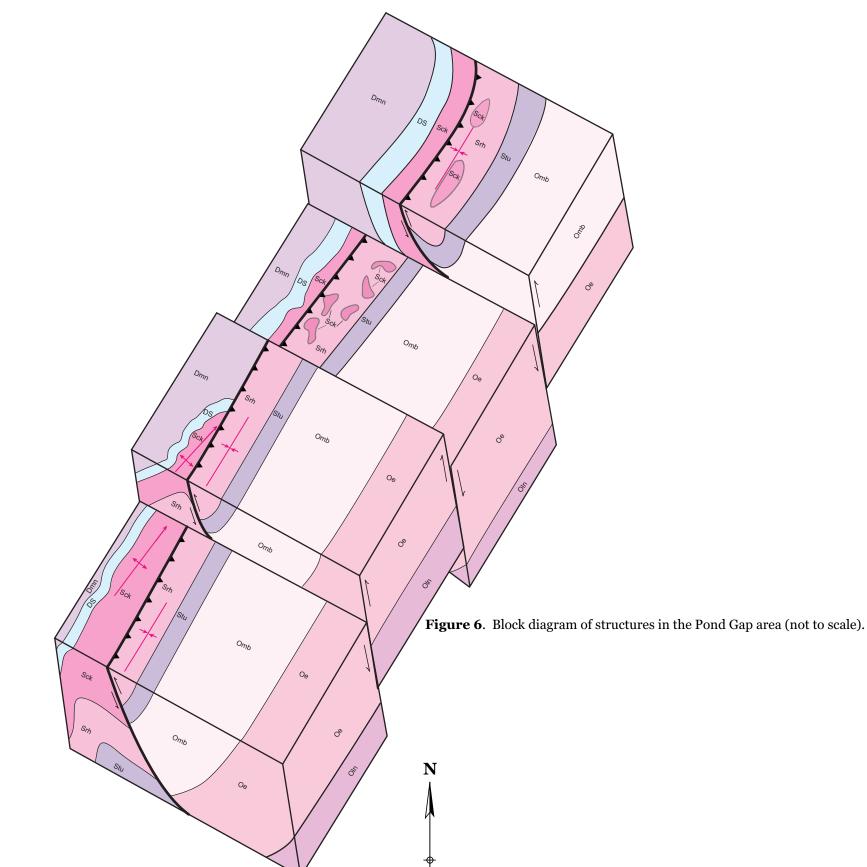
2. Subsurface structures interpreted from surface measurements

NORTH

MOUNTAIN FAULT

Foliation-showing trend and plunge of intersection

Direction and angle of plunge of minor anticline



North Mountain fault - The North Mountain fault is a regional thrust fault that dips 20 degrees or less to the southeast within the quadrangle. Thrusting of hanging wall rock of the Cambrian Waynesboro Formation over the Ordovician Beekmantown Formation along the fault resulted in thousands of feet of displacement. Fault breccia and highly fractured rock occur in a belt up to hundreds of feet wide along the trace of the fault. The breccia can contain fragments of country rock of varying sizes (Figure 4) or have the appearance of crush conglomerate.

STRUCTURAL GEOLOGY

Little North Mountain fault - To the north of the quadrangle, the North Mountain fault splits into two fault traces. The North Mountain fault proper swings east into the Great Valley while a second fault continues south, along the west flank of Little North Mountain. This fault is herein referred to as the Little North Mountain fault. In the Augusta Springs quadrangle, this fault places Ordovician and Silurian age rock over Silurian and Devonian age rock. The Little North Mountain fault dips approximately 30 degrees or more to the southeast and has an apparent displacement of several hundred feet. Brittle deformation is evidenced by the angular fragments of brecciated rock along the fault (Figure 5). There are at least three cross faults that displace the main fault along its trace through the quadrangle. These cross faults are high-angle faults that create discrete fault blocks between them (Figure 6).

Estaline Valley syncline - Estaline Valley is underlain by Devonian Millboro Shale that is roughly in a synclinal form. The Millboro is deformed by interformational folding and faulting (Figure 7). There is one small amplitude anticline near the eastern edge of the synform that brings the Ridgeley and Helderburg to the surface.

Brown Ridge anticline - Brown Ridge is formed by an anticline. Although there does not appear to be any section missing, the contact of the Ridgeley Sandstone and the Needmore/Millboro map unit is convoluted and the Helderburg Group has been slightly metamorphosed. The anticline plunges northeast into deformed Millboro Shale.

Elliott Knob syncline - Only a small part of the east limb of the Elliott Knob syncline is in the quadrangle. Devonian rocks dip to the northwest with minor folding and faulting (Figure 8).

ECONOMIC GEOLOGY

Limestone and dolostone - Cambrian, Ordovician, and Devonian carbonate rocks crop out in the quadrangle and have been sampled and analyzed by Giannini and Hostettler (1994). There are no active large quarries in the quadrangle, but there are small quarries that were made for local use. Potential use for these rocks include aggregate, dimension stone, and high-calcium limestone.

Shale - Devonian shales have been tested in the quadrangle for structural clay products and have been found suitable for brick and lightweight aggregate (Calver, Smith, and LeVan, 1964). An abandoned shale pit near Fordwick was operated by Lehigh Portland Cement Company in the 1950's and 60's for use in the manufacture of portland and masonry cement. One shale pit is currently intermittently active for use as road ballast within the state wildlife management area.

The ore is probably associated with the Ridgeley Sandstone; it may have formed as a secondary deposit along a fault. Concentration of minable ore is limited to small areas and does not appear to be regional in extent. The ore was utilized in iron making at three furnaces in the quadrangle: Ferrol, Estaline, and a reported furnace at Pond Gap (now Augusta Springs).

Sandstone and quartzite - The Tuscarora and Keefer sandstones and possibly

Iron - Iron ore was mined from small open pits in the late 1800's (Figure 9).

Testing for these commodities has not been done; however, there are several promising exposures within the quadrangle that may have the potential. **Bauxite** - Prospect pits for bauxite were dug in the early 1900's on the Yago farm. Several pits, collectively called the Yago prospect, are quite large (Figure 10), and all occur within the Beekmantown Formation. There is no indication of

the Ridgeley have potential for high-silica sand and non-polishing aggregate.

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SUGGESTED REFERENCE

Portions of this publication may be quoted if credit is given to the Virginia Department of Mines, Minerals and Energy, Division of Geology and Mineral Resources. It is recommended that the reference to this report be made in the following form:

Coiner, L. V. and Wilkes, G. P., 2008, Bedrock geologic map of the Augusta Springs quadrangle, Virginia: Virginia Division of Geology and Mineral Resources, Open-File Report 08-01, 1:24,000- scale map.



Figure 1. Augusta White Sulphur Spring issues from the Licking Creek Limestone.



Figure 2. Near vertical ledge of Tuscarora sandstone on Little North Mountain. Because of the Tuscarora's resistance to weathering, it underlies the crest of the mountain in many places.



Nodules are diagnostic of the Lincolnshire.



Figure 4. North Mountain fault breccia. The rock consists of varying sized clasts of country rock and calcite-filled fractures.



Figure 5. Little North Mountain fault breccia. Angular to sub-angular clasts of Silurian sandstone indicate the brittle nature of the faulted rock.



Figure 7. Northwest-verging chevron fold in an outcrop of Millboro Shale on Route 601 in the Estaline Valley southeast of the Brown Ridge fault.

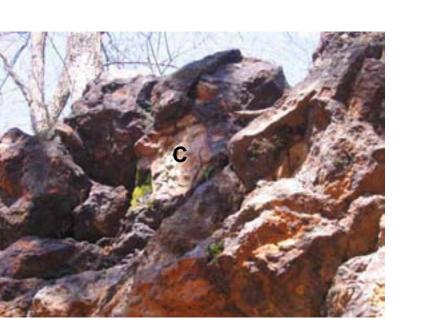


Figure 9. Iron ore outcrop in mine near the Pond Gap Road. The large angular clast of country rock (C) that is surrounded by limonite is about 1 foot across. This and other iron mines contained ore that was pod-like and was extracted by open pit methods. The ore was processed into pig iron at nearby furnaces.



Figure 10. Largest of the Yago farm prospect pits. Prospectors sought aluminum ore in the Great Valley during the 1940's.

Mineral Resources - identification numbers are preceded by "177-C" in the Mineral Resources of Virginia database.

Prospect (sh=shale, fe=iron, al=aluminum)

Active mine or quarry (ls=limestone, sh= shale)

Abandoned mine or quarry (ls=limestone, fe= iron,

sn= shale)

Geologic References

R-1001 Rock repository sample
 R-1200(G) Rock repository sample with geochemical analysis
 C-3000 Other geochemical analysis (C-carbonate)
 Location and figure number of photograph



Figure 8. Deformed Devonian-Silurian rock within the Estaline Valley syncline. Field of view is approximately 8-by-15 feet.